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***Glacial geology and  
geomorphology of  
Weardale.***

***Volume II***

***by***

***Edwin Neville Moore***

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***A thesis submitted for the degree of  
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April, 1994.***



**27 JUN 1994**

Weardale 1:10,560 series

Field Map Extracts:-

MAPS	1.	NZ.13.SW	Bedburn confluence South
	2.	NZ.13.SW	Harperley
	3.	NZ.13.SW	Bedburn confluence North
	4.	NZ.13.NW	Bradley
	5.	NZ.03.NE	Wolsingham East
	6.	NZ.03.NE	Wolsingham West
	7.	NA.03.NW	Frosterley
	8.	NZ.03.NW	Rogerley
	9.	NY.93.NE	Stanhope to Eastgate
	10.	NY.93.NW	Eastgate to Brotherlee
	11.	NY.93.NW	Brotherlee to Westgate
	12.	NY.83.NE	St John's Chapel
	13.	NY.83.NE	Wearhead

# Glacial Geology and Geomorphology of Weardale

## Volume II

### Field Techniques

This volume includes a selection of the original field maps completed during the survey. The sequence of maps included embrace the main terraces and associated morphological features of the Dale. These features have been mapped on the O.S., 1:10,560 series. Detail was later transferred to O.S. maps, 1:25,000 series (Vol.1., Figs. 10.1 - 10.11). The area originally covered embraced all sheets NY84/94, NY83/93, and N213.

#### 1. Aims and Objectives in the field.

The primary aim of the study was to assess the extent and nature of Devensian glacial process in the Dale. This necessitated detailed morphological mapping of the area covered by Pennine Till. This cover extended from the head of the Dale at Killhope Cross (GR.799.433) to the Scars (GR.136.319), below the Bedburn confluence at Witton-le-Wear, a distance of some 42 kilometres. No erratics had previously been found in any remnant of till on interfluves around the Dale, to the west of High Stoop, Tow Law (GR.164.403) in the north, or of Five Pikes (Gr.013.369) in the south.

The current survey acquired access to many exposures of till which cap the interfluves above Coal Measures strata. To the north of the Dale this strata forms the slopes and interfluves to the west of the Waskerley Beck. South of the Wear, Coal Measures cap the interfluves from the Ayhope to the Harthope and the Wear. These beds finally form the underlying strata to the east of the Bedburn.

Exposures in till confirmed a presence of Pennine ice at Inkerman, Tow Law. East of this location, at Stonefoot Hill, at Sunnyside, and Dowfold Hill, Crook, tills included erratics which had been derived from Lake District and Scottish sources. South of the Dale, erratics of Lake District and Vale of Eden origin were also found in tills to the east of the Bedburn drainage.

These exposures, along with examination of glaciofluvial deposits, confirmed the area of the Dale around Witton-le-Wear as being one where ice flows had merged.

Mapping of the Dale to the west of Witton-le-Wear embraced an area of some 630 sq. km. At Greenfoot, Stanhope the interfluves are at a distance apart of some 14 kilometres (i.e. from Horseshoe Hill, (GR.985.499) in the north, to Raven Seat, (GR.970.233) on Harnisha Hill to the south).

The extent of this area of field study necessitated locating a base within the Dale. The original mapping was

completed using White House Farm, Eastgate, as this base. (thanks are recorded to Mr and Mrs W. Rippon who allowed siting of a caravan at this location). Later studies were completed after taking up residence in the Dale at Witton-le-Wear.

## 2. Mapping

Mapping embraced a study of the river terraces, glaciofluvial forms, till morphology, as well as meltwater channels left by deglaciation. The extent of the study area necessitated an aerial photographic study of the Dale. The University of Durham stereoscopic aerial photographic cover of the Dale proved invaluable as a source of supplementary evidence, both for, and after, the acquisition of field data (eg. The Greenly Hills moraine - see Vol. 1. Ch. 5). Stereoscopic survey also aided identification of morphological features which were concealed by topography, or vegetative cover at ground level. Features identified on these photographs were mapped subsequent to a ground survey which included a geological survey of surficial deposits and solid strata.

Levelling of features e.g. meltwater channels and paired terraces was carried out with the use of a quick-set level. Slopes which were difficult of access (e.g. The gorge at Greenly Hills moraine (see Vol. I. Ch. 5)) were measured by clinometer. An aneroid checked at hourly intervals to

known spot heights, was utilised for recording heights of features on higher fells above 450 m O.D.

Meltwater channels were mapped which would include those formed by marginal and sub-marginal flow, as well as larger ice-directed channels, (cf. Mannerfelt, 1945, 1949; Sissons, 1958, 1960a, 1960b, 1961a; Derbyshire, 1961; Clapperton, 1968; Price, 1973). Channels are listed according to location, height of incision and gradient of floor (see Vol. 1 Ch.7).

### 3. Further Objectives

A Quaternary field survey necessitated defining the following objectives:-

- a) An assessment of the quality of solid strata in the Dale and to what extent the degree of weathering was derived from Quaternary process. Exposures of solid strata were fractured to provide clean sections for examination of rock quality, and for the presence of open, or iron-stained joints.

These observations were later correlated with details of the quality of solid strata encountered by exploratory boring for the Kielder Tunnels Project.

- b) A Fabric analysis of all exposures of basal till.

Clasts in the diamict were hand sectioned for field identification. Where further identification was required, thin sections of the rocks were prepared, then analysed with a petrological microscope (cf. Bridgland, 1986). A minimum of fifty clasts were examined at each till exposure. A total count of only 50 gives a standard error of some 13.9% (Bridgland, 1986). However such counts may still be considered viable if correlated with data from neighbouring locations. Records were taken of parent lithology, size and 'a' axis orientation. Similar records were taken of large gritstone boulders (1-3 m across) deposited in scattered trains across several interfluves to the Dale.

- c) Recording of sedimentary content and structures in glaciofluvial deposits. Detail of Weardale glacial sediments (at depth) was provided by access to borehole data from the Blue Circle factory, and the Kielder Tunnels project. (kindly see acknowledgements Vol.1). Road construction also provided detailed sedimentology of glacial kames, terrace gravels, and sections through tills and solifluction deposits (e.g. Willow Green Ch. 5). Excavation for bridge pillars at Broadwood, Frosterley and at the A.68 Bypass Witton-le-Wear provided further details of sand and gravel sequences in the Quaternary terraces. Details of the glaciofluvial content of kames was also provided by

extraction of these deposits for track surfacing.

Sampling of slope deposits for evidence of mass movement found inclusions of these deposits within the older gravel terraces of the Dale.

Consequently older terrace gravels are defined as those which include interdigitated, and thus contemporary with, soliflucted deposits from adjacent slopes. Borehole details reveal their sedimentology to also include ablation till, boulders, and silts suggesting derivation from valley sandurs. Artifacts included on their surface suggest these terraces formed dry sites by Mesolithic times when they became utilised by man.

Incision of the older gravels occurred with rapid runoff as a result of deforestation in Iron Age/Romano-British times. Lower terrace levels so formed are capped by silt. Reworking of former gravels also occurred with deposition of a terrace of mineral rich sediments, notably fluorspar, as a result of anthropogenic 'flash flooding'. This resulted from the North Pennine orefield practice of hydrological mining termed 'hushing'. This practice was very active during the operations of the London Lead Company 1762-1844.

d) An evaluation of the nature and effects of periglacial processes during the Quaternary in the Dale.

Deposits which include masses of sharp angular fragmented debris orientated downslope were assigned to a derivation by frost process. These were found in localities where exposures of strata, and inclined boring for the Kielder Tunnels, revealed clay-filled open joints. Such joints were also examined in detail thanks to access to the experimental tunnel at Rogerley Quarry (kindly see acknowledgements). Joints were also examined at valley quarries and at surface exposures.

Paraglacial processes were also operative in the Dale. These appear to have provided the spreads of slumped deposits from reworking of the basal till. Such deposits are reddish sandy clays. In contrast to the blue-grey basal tills they lack an argillaceous content and show evidence of oxidation and deep weathering.

Patterned ground is of limited occurrence in the Dale. Sampling occurred at the limited localities.

Lithology, size and orientation of block fabric in Felsenmeer were recorded. Orientation of 'a' and 'b' axes were taken by prismatic compass. Dips, where

applicable, were taken with a clinometer.

e) Recording of glaciotectonic structures.

Glacial geological investigations encountered evidence for glaciotectonic disturbance of strata. Greenly Hills form an apparent push moraine (see Ch.5). Extensive glacial rafting was also found in Westphalian strata, both within the Wear valley, and across both interfluves to the Dale. The presence or absence of permafrost in the Westphalian beds appears to have contributed to the extent of strata rafted, and determined the degree of disturbance at the location of extraction, as well as in the included beds. Analysis of structures on the interfluve, necessitated comparisons with other glaciotectonic disturbances observed in similar beds at locations north of the interfluve. At the latter permafrost appears to have been absent (see Vol. 1. Ch.6). Thanks are recorded for access to sites where the rate of open-cast coal extraction can result in several hectares being extracted and infilled within one week.

4. Analysis of the Data.

Analysis of the data has been presented by:-

1. Tabulation of data recording the quality of solid rock strata.

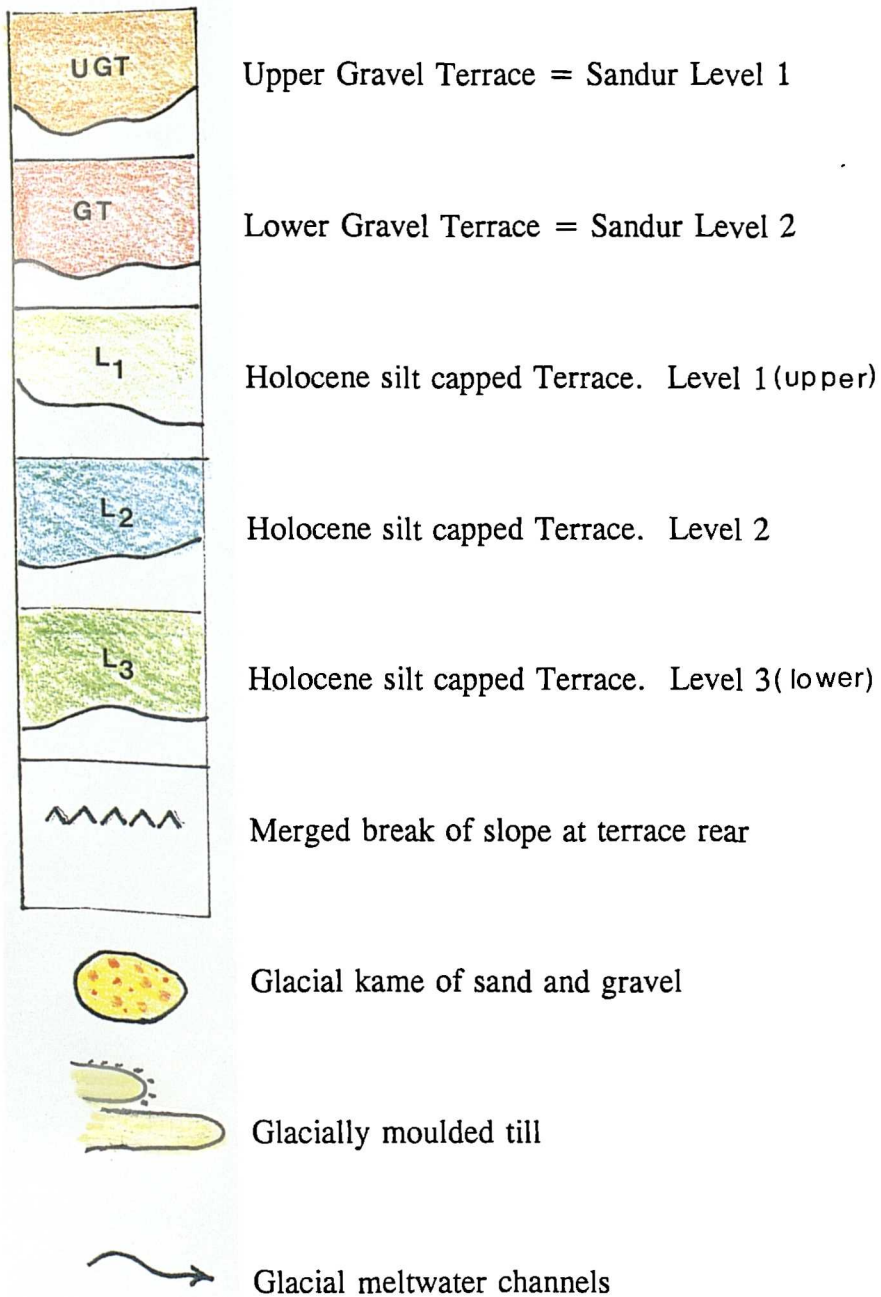
2. Plotting of polar scattergrams and modal azimuths of stone content in basal till.
3. Plotting of polar scattergrams and modal azimuths of blocks scattered across interfluves and in Felsenmeer of the Dale.
4. Geological sections showing the nature of till, glaciofluvial, and sandur deposits of the Dale.
5. Geological sections of glaciotectonic rafts.
6. Tabulation of meltwater channels based on location, height, average gradient, and incision.
7. Plotting of all data onto a geomorphological map of the Dale. Details are also recorded on the O.S. 1:25,000 series (Vol.1 10.1-10).

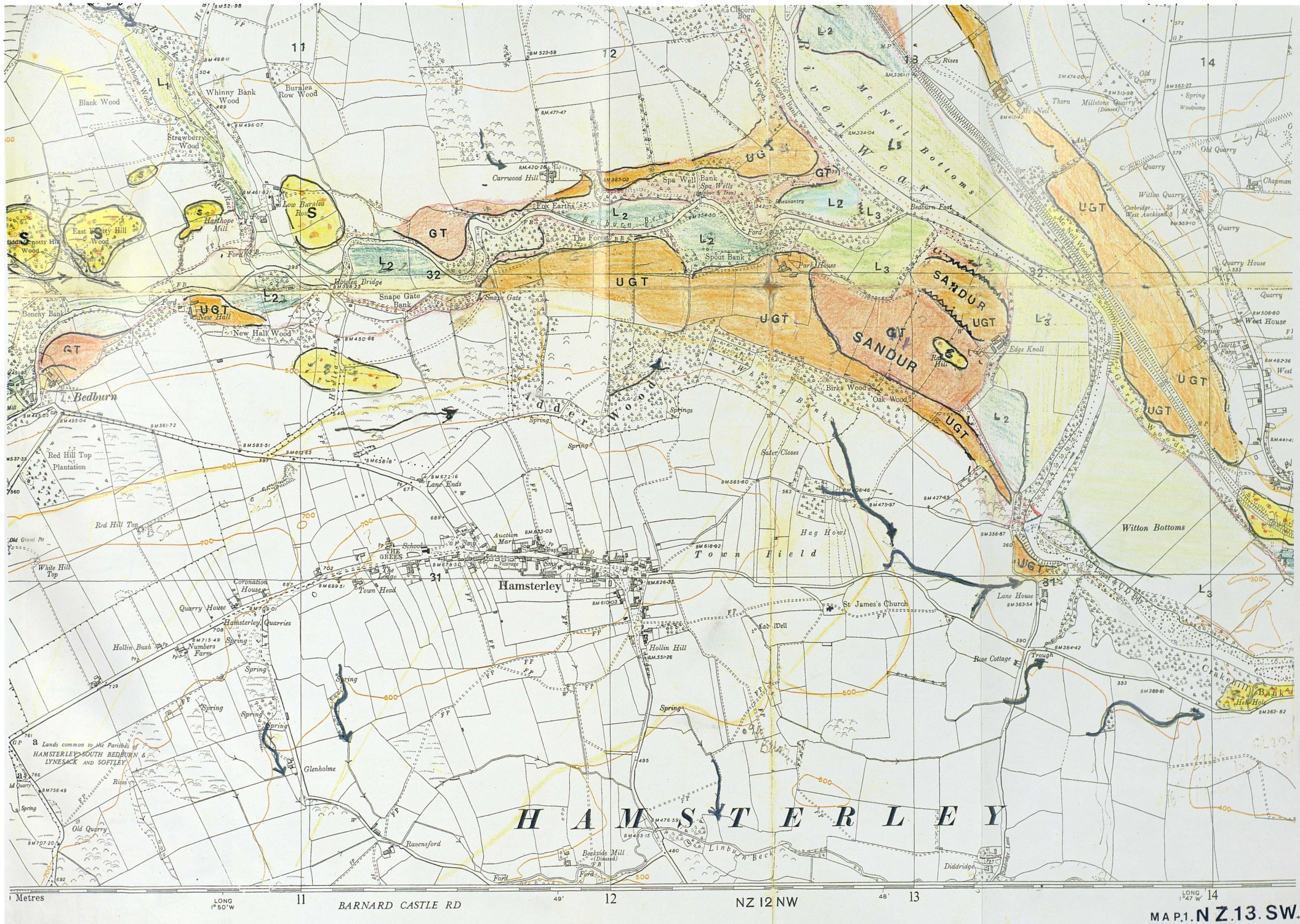
The analysis is based on an acquisition of field data. Such data lacks an accuracy of refinement by further laboratory studies. However the correlation of that data presents a hypothesis which it is hoped may serve as a challenge to further detailed explorations.

Extracts from the original field maps are to be found in the subsequent pages.

Key to 1:10,560 Field Maps

Grid Squares. 1km Square.





1 Metres LONG 1° 50' W 11 BARNARD CASTLE RD 49' 12 NZ 12 NW 48' 13 LONG 1° 47' W 14 MAP.1. N Z.13. SW.

1000 500 0

80

DURHAM

MAP. 2. NZ. 13. SW.

WEARDALE RD

WOLSHINGHAM PH 13

NORTH WEST DURHAM CO CONST

LONG 1° 47' W

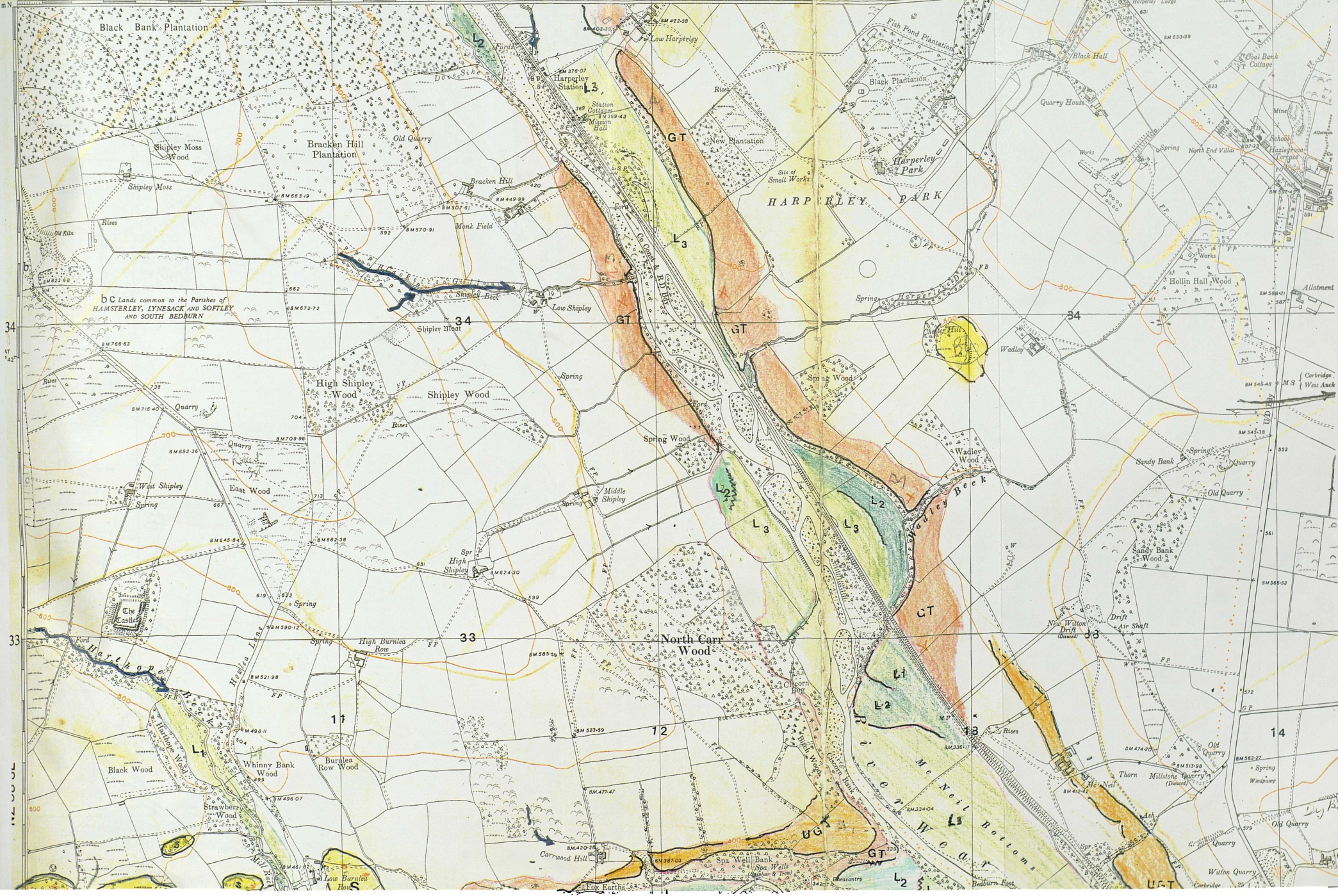
11

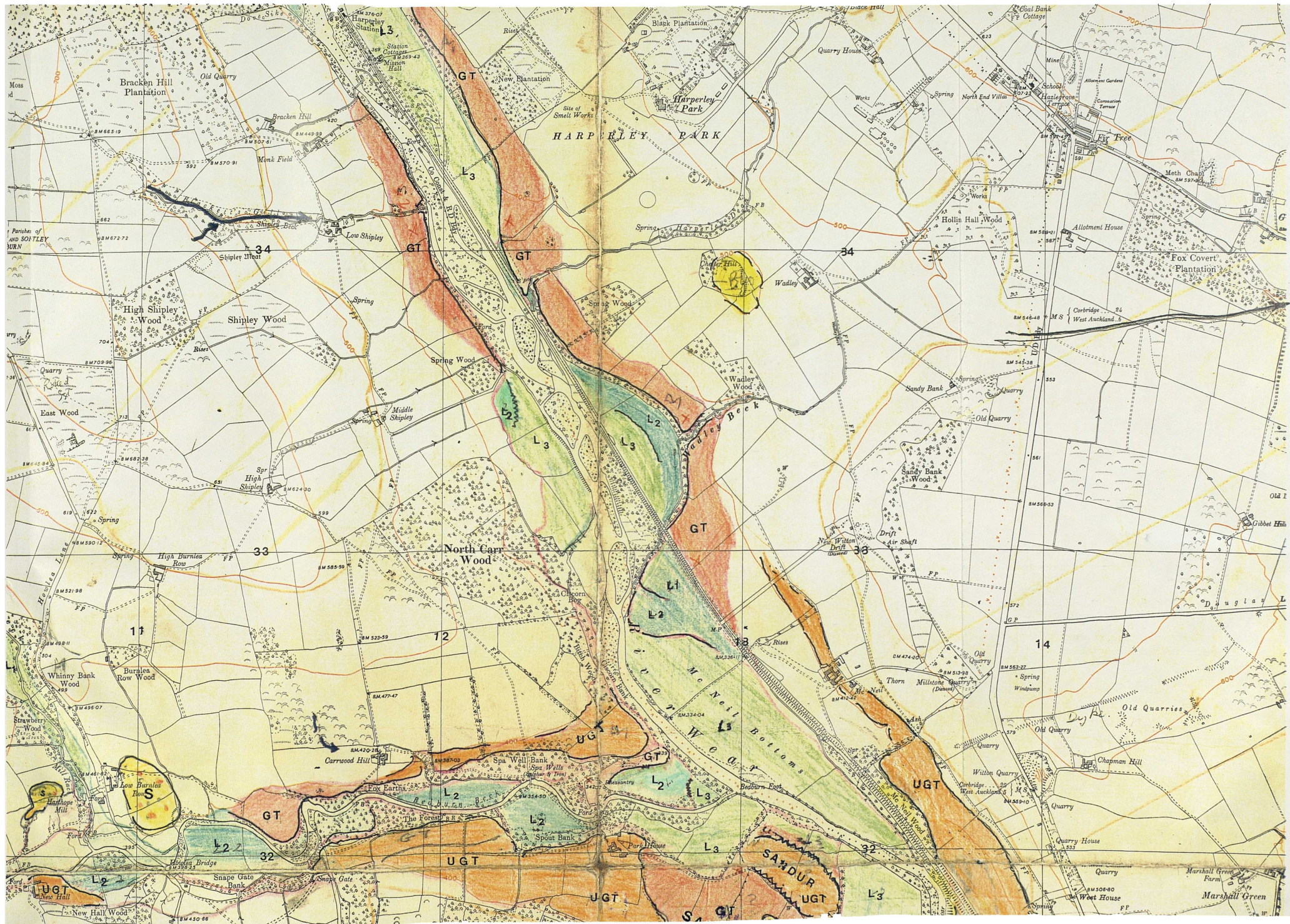
12

NZ 13 NW

14

1000 Metres







Black Bank Plantation

Final Edition on National Sheet Lines)

*a flow of side*

DURHAM

NZ 04 SE 63

160

WEARDALE RD 52

08

09

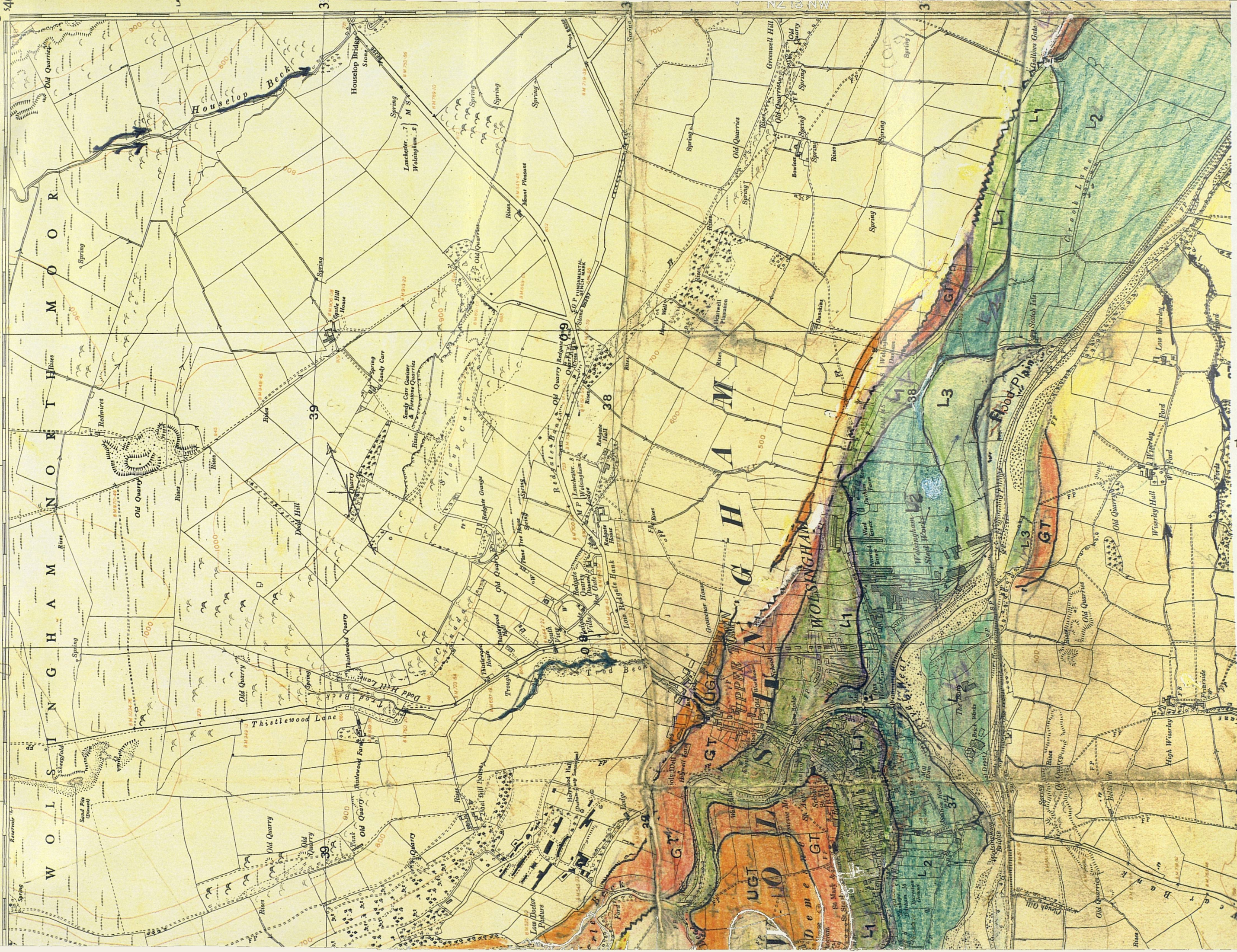
09

240 C

10000

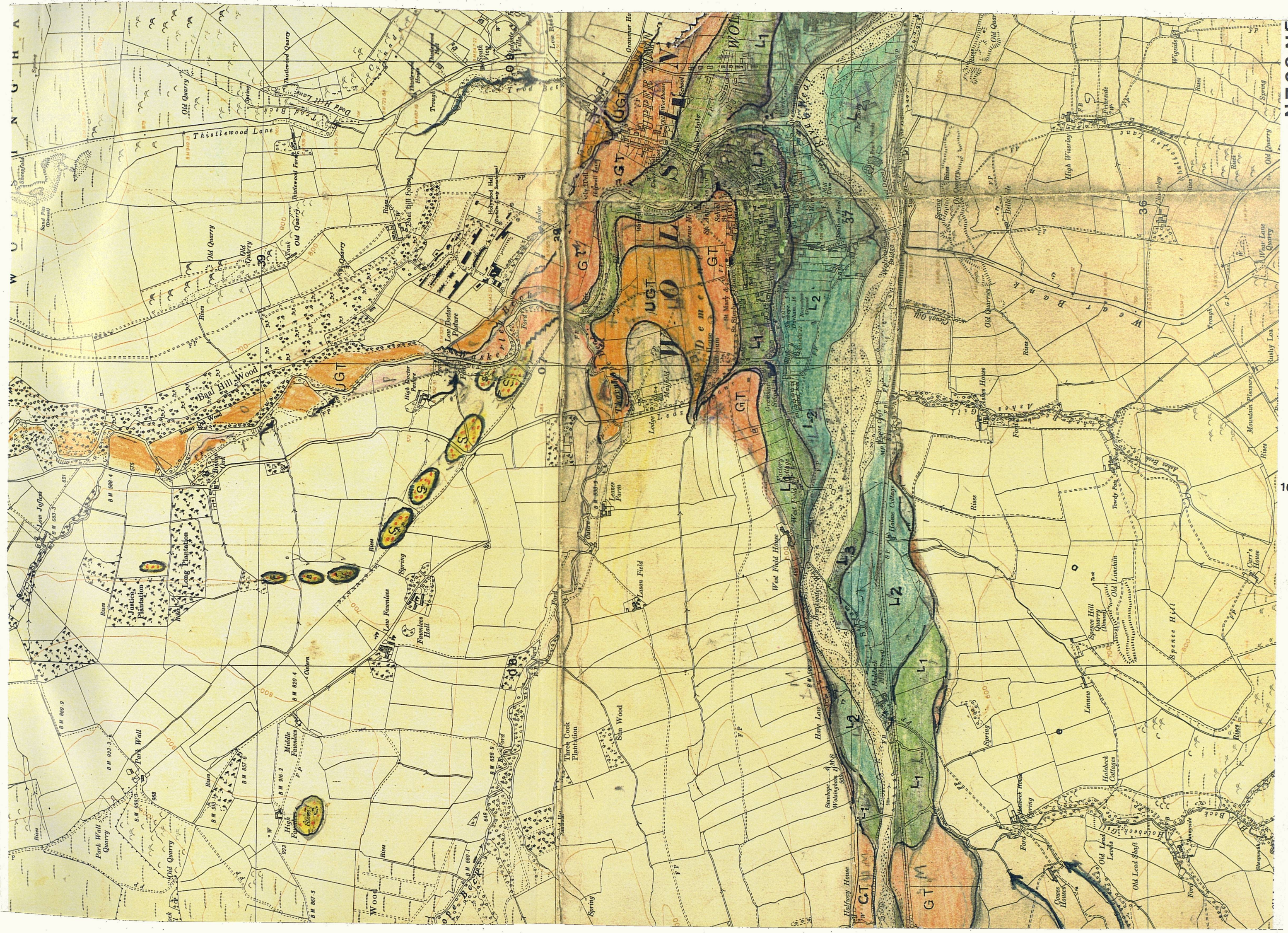
LONG 1° 1' W

340



15

MAP. 5. N Z. 03. NE.







96

97

98

99

Bewdley Plain

Green Head

STANHOPE

UGT

UGT

UGT

96

97

98

99

Long Plantation

Hag Top

High Horsley

Old Shaft (Lead)

Old Limekiln

Old Limekiln

Old Shaft (Lead)

Quarry Hill

Hothill Quarry (Diamond)

Horsley Head

Old Quarry

Old Shaft (Lead)

High House

Shield Ash

Moorhouse Side

Old Clay Pit

MAP 9. NY.93.NE.

19









North 1117  
Scale 1:20,000

# HART HOPE MOOR

Pearl Riggs

MAP. 13. NY. 83. NE